## GEOTHERMAL MAPPING USING TEMPERATURE MEASUREMENTS

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## **THE EARTH'S TEMPERATURES**

- Earth's temperatures remain relatively constant over time except in the very near surface
- The temperature difference between the cool crust and hot molten magma causes a temperature gradient
- Normal temp gradient in deep boreholes around the world shows an increase of 15 to 30°C per km.
- In geothermally active region temperature increases more rapidly with depth, hence, higher heatflow.



#### The Earth's Temperatures, cont'd

- For high temp resources development, there's need to target areas with elevated temp gradient.
   How? (heat-flow measurements)
- For low temp resources development, the relatively stable earth's temp is sufficient (e.g. geothermal heat pumps).



#### **THEORY OF HEATLOSS SURVEY**

- Heat loss analysis involves evaluation of heat transfer in a geothermal system.
- Heat transfer mechanism is divided into three major processes
  - conduction
  - convection.
  - Radiation



## CONDUCTION

Occurs when;

- When adjacent atoms vibrate against one another, or as electrons move from atom to atom.
- No flow of the material medium.
- Occurs mainly in solids.



## **Conduction**, cont'd

Conduction heat flow is calculated by using one dimension heat conduction equation

$$Q = Ak \frac{dT}{dy}$$

#### Where

- Q Conductive heat flow (watts),
- A Surface area of hot ground (m2),
- k = 2 Thermal conductivity of rock (w/m°C),
- T Temperature (°C)
  - Depth (m).



### Sample calculations for Conductive heat transfer

	East (Km)	North (Km)	T <sub>s</sub> (°C)	T <sub>50</sub> (°C)	T <sub>100</sub> (°C)	Grad <sub>50</sub> (°C/m)	Grad <sub>100</sub> (°C/m)	Average (°C/m)	Mean (ºC/m)	Area (m2)	Heat flow (MW <sub>t</sub> )	
H-1	190.9	91.0	24.3	28.4	28.4	4.1	4.1	4.1	14.6	5000	73	
H-2	188.4	93.5	95	95	95	0	0	0				
H-3	188.3	93.5	28	34	36.2	6	8.2	7.1				
H-4	187.9	93.4	61.3	97	97	35.7	35.7	35.7				
H-5	187.6	93.2	35.4	57	68	21.6	32.6	27.1				
H-6	186.8	83.2	23.1	28.8	29.7	5.7	6.6	6.15				
H-7	189.3	86.7	24	30.7	32.4	6.7	8.4	7.55			3	)
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#### **CONVECTIVE HEAT TRANSFER**

Convective heat transfer occurs in hot springs and fumaroles

involve transfer of heat energy due to movement of fluid particles.

This process is faster than conduction.



## **Convection, cont'd...(2)**

• Calculation of convective heat transfer is calculated using the fluid flow equation

$$V = C_d A_t \sqrt{\frac{2g\Delta H(\frac{\rho_w}{\rho_s} - 1)}{(1 - \frac{d_t^2}{d^2})}}$$

#### Where

- V Volumetric flow rate (m3/s),
- Cd Coefficient of discharge (assumed to be 0.96),
- g Acceleration due to gravity (9.81 m/s2),
- dt Venturimeter throat diameter,
- d Venturimeter diameter at the high pressure tapping (m),
- $\Delta H$  Differential height at the manometer (m),
- $\rho$  Density of the fluid (kgm-3),
- $A_t$  Throat area (m2)



## Convection, cont'd...

• Equation below is used to calculate the convective heat flow.

$$Q_c = V \rho_s h$$

- Where
  - Qc Convective heat flow (watts)
  - h Enthalpy of steam at the corresponding measurement temperature (Jkg-1).



## Sample calculations for convective heat transfer

PROPERTIES	SYMBOL	VALUE	]
Venturi diameter	d	0.0254 m	
Water density	$\rho_w$	1000 kg/m <sup>3</sup>	
Diameter (throat)	d <sub>t</sub>	0.00635 m	
Steam density	$\rho_s$	0.4753 kg/m <sup>3</sup>	
Throat area	A <sub>t</sub>	3.16532E <sup>-05</sup> m <sup>2</sup>	
Steam enthalpy	h <sub>s</sub>	2670 kJ/kg	
Coef. of discharge	Cd	0.96	
Water enthalpy at ambient conditions	h <sub>o</sub>	117 kJ/kg	
Gravity	g	9.81 m/s <sup>2</sup>	
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## Sample calculations, cont'd

Name	North (km)	East (km)	Man. height (mm)	Volumetric flow (m <sup>3</sup> /s)	Heat flow (kW <sub>t</sub> )
KFMRL- 1	177.0	89.2	9	0.000604765	0.767
			17	0.000831169	1.055
			9	0.000604765	0.767
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## **Temperature and Pressure in Boreholes**

#### • Temperature measurement

- Boreholes serve as gradient holes for temp gradient computation.
- Results used to estimate geothermal reservoir temperatures

#### Pressure measurements

- Assist in modelling the hydrological picture of the area
- Give an indication of possible recharge and outflow zones.



## **Tools Used during heat-flow survey**

- GPS
- Digital Thermometer
- Fabricated spikes
- Hammer
- Field note book
- Metal rod
- Field Map
- Winch
- Kuster temperature and pressure tools
- Jembe, Spade, Plastic basins, Panga
- Reliable means of communication e.g radio calls, Satellite phones, mobile phones etc



## PLANNING FOR HEAT-LOSS SURVEY

- Spacing of gradient holes
  - 100 m 1 km in an area of high thermal activity
  - 1 4 km in an area of low activity.
- Area, time and resources available

   This determines how fast the work is to be accomplished
- Geological formation of the area
   –Some important features
- Social constraints

-Areas prohibited or controlled (religious, tourist attraction, cultural)



# Heat-loss survey and other Geo-scientific studies

- Assists in quantifying amount of heat being lost on the surface
- Complements other disciplines in:
  - determining the reservoir temperature
  - identification of active structures
- Gives extent of leakage through the capping.
- suggests possible orientation of the fracture zones



## **Challenges Encountered**

- Hostile climate
- Poor communication (no roads)
- No proper tools and equipments
- Some areas are covered with hard material which are difficult to penetrate
- No surface manifestation so convective heat transfer becomes difficult to measure
- Heavy and cumbersome equipments



#### **CASE STUDIES (EXAMPLES IN KENYA)**

Heatflow measurement has been carried out in five of geothermal prospects

- Menengai-Olbanita Prospect (2004)
- Baringo Prospect (2004)
- Arus-Bogoria Prospect (2005)
- Korosi and Chepchuk Prospect (2006)
- Paka Prospect (on-going) (2006 on-going)



There are 14 geothermal prospects in Kenya

Total estimated geothermal potential >3000 MWe



## **Menengai-Olbanita Prospect (2004)**

- The prospect is associated with the 90 km<sup>2</sup> Menengai Caldera
- Collection of heat flow data was done in an area of about 900 km<sup>2</sup>.
- No hot springs were encountered.
- Convective heat flow was obtained from the steaming grounds within the caldera
- Boreholes drilled to 300 m depths in this area discharge water at temperatures of 40 °C to 60 °C



## Menengai Caldera





## Menengai-Olbanita Prospect, cont'd

- orientation of the high temperature features are – NNW-SSE, and NE-SW
  - major fault/fracture are also in this direction
- Over 3536 MWt heat lost naturally.
   2690 MWt is lost in the Menengai Caldera
- This large heat loss could be an indicator of a huge heat source underneath this prospect



## **Menengai-Distribution of Gradient holes**



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## **BARINGO PROSPECT (2004)**

- Heat source is due to dyke swamps along fault lines
- orientation of the high temperature areas is NE-SW
- Coincides with those of major fault/fracture zones.
- Total heat loss from the prospect is > 1049 MWt
- Conduction
  - 941 MWt
  - 90% of the conductive heat loss occurs along the fault zones

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- convection
  - 108 MWt by (105 MWt is lost in Kokwa Island)

## **Baringo-Distribution of Gradient holes**



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## Arus-Bogoria Prospect, cont'd

- Total heat loss from the two prospects is in excess of 1666 MWt
- Conductive 1229 MWt

   Lake Bogoria 762 MWt
   Arus 467 MWt
- Convective 437 MWt

   Lake Bogoria 437 MWt
   Arus 0.03 MWt





#### Geyser at Lake Baringo

#### Arus Steam jets at Arus



## **Arus Bogoria-Distribution of Gradient holes**





## **KOROSI AND CHEPCHUK PROSPECT**

- Heat source at Korosi is therefore controlled by NE-SW and NW-SE trending faults.
- Conduction
  - About 2,135 MW<sub>t</sub> Korosi
  - About 546 MW at Chepchuk.

Convection

- 0.4 kW<sub>t</sub>
- Almost all the heat lost is by conduction





#### **Korosi Prospect**







## **PAKA PROSPECT**

- Paka is a caldera north of Korosi
- Surface manifestations cover about 45 km<sup>2</sup> (hot grounds and fumaroles)
- Surface studies are in progress and expected to end in January, 2007







Paka Prospect

## **Comparison between prospects**

Prospect	Conduction (MWe)	Convective (MWe)	Total (MWe)
Menengai- Olbanita	1060	2476	3536
L. Baringo	941	108	1040
Arus- Bogoria	1229	437	1666
Korosi- Chepchuk	2681	0.4	<b>2681</b>
Paka	<b>On-going</b>		



## CONCLUSION

Heatloss survey is an important tool in;

- Analysis of the distribution of the heat loss features
- Results obtained are used as an indicator of the heat source size
- Give an indication of the magnitude of recharge
- extent of leakage through the capping.
- determining the reservoir temperature
- Serve as a guide in locating hidden fracture zones
- Ranking the prospect for development



## **THANK YOU**

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